EAS270, "The Atmosphere"

Assignment 1

Due 5 pm 19 Oct., 2011

<u>Professor</u>: J.D. Wilson

<u>Value</u>: 10%

**Instructions**: All questions have equal value. Retain high precision in your calculations, but please round your final answers to three significant digits<sup>1</sup> and state the unit (examples:  $20.1^{\circ}$ C,  $1.09 \text{ kg m}^{-3}$ ,  $391 \text{ W m}^{-2}$ ,  $4.11 \times 10^2 \text{ W m}^{-2}$ ). Document your working tidily — there is potentially a one mark penalty for illegible or unintelligible working. Please drop off your assignment (labelled with your name and ID number) in class, or in the drop-off box outside Tory 3-40. A two mark penalty will be applied for late assignments received before noon Thursday 20 Oct. After that time, the late penalty will be five marks.

**Task**: Add together the *last five digits* of your student I.D. number, to form what we shall interpret as a ground-level temperature  $T_1$ , as a latitude  $\phi_1$ , and as a surface shortwave reflectivity (albedo)  $\alpha_1$  expressed as a percentage, e.g.

I.D. number 1198765  $\rightarrow T_1 = 35^{\circ}$ C,  $\phi_1 = 35^{\circ}$ ,  $\alpha_1 = 35\% \equiv 0.35$ .

In what follows, the subscript "1" denotes properties at level 1, which is ground level. Level 2 (denoted with subscript "2") will be the 850 hPa level.

- 1. Assuming the emissivity of the ground surface is  $\epsilon = 0.96$ , compute the emitted longwave radiative flux density " $L_1 \uparrow$ " corresponding to your ground temperature  $T_1$ .
- 2. Compute the ratio of this flux  $L_1 \uparrow$  to the flux " $L_{\text{ref}} \uparrow$ " that would be emitted by the same surface if it had a temperature of 0°C (273.15 K).
- 3. Based on your  $L_1 \uparrow$  but neglecting incoming (i.e. downward) longwave radiation  $(L \downarrow)$ , compute the net radiation " $Q_{*1}$ " at solar noon at the times of the equinox and express this as a fraction of the solar constant  $Q_{*1}/S_0$ . (To compute  $K \downarrow$  assume the solar beam is not subject to scattering or absorption by the atmosphere, i.e. is transmitted without attenuation so that  $K \downarrow$  depends only on the latitude  $\phi_1$ .)
- 4. Your net radiative energy supply  $Q_{*1}$  would suffice to evaporate what depth  $d_1$  of water over a period of one hour? (Take the density of liquid water to be  $\rho_w = 1000 \text{ kg m}^{-3}$ )
- 5. Assuming ground-level pressure is  $P_1 = 925$  hPa, compute the ground-level air density  $\rho_1$  implied by this combination  $(P_1, T_1)$ .
- 6. Assuming the specific humidity at ground is  $q_1 = 0.002$ , compute the vapour pressure  $e_1$ .
- 7. Compute the mass of water  $m_1$  in one cubic metre of air at ground level.

<sup>&</sup>lt;sup>1</sup>However if you are providing an answer in Kelvin units, it is appropriate to provide five significant figures, e.g. 278.75 K.

- 8. Adopt the hydrostatic law (given as data), evaluating the right hand side as  $-\rho_1 g$ , to compute the vertical distance  $\Delta z$  between the ground and the 850 hPa surface (i.e. the altitude where pressure is  $P_2 = 850$  hPa). Note: here we are using the hydrostatic law in an approximate way, by placing a fixed constant value of  $\rho (= \rho_1)$  on the right hand side.
- 9. Assuming the temperature variation from ground to the  $P_2$  level follows the dry adiabatic lapse rate (DALR), compute the temperature  $T_2$  (in Celcius units) at this level.
- 10. Again use the hydrostatic law, evaluating the right hand side as  $-\rho_2 g$ , to determine the pressure  $P_3$  at a point that is 300 m above the 850 hPa surface.

## Data

- 1 hPa = 100 Pa,  $T [K] = T [^{\circ}C] + 273.15$  (Note that a *change* of one degree Kelvin is the same as a *change* of one degree Celcius).
- $\frac{\Delta P}{\Delta z} = -\rho g$

The hydrostatic law.  $\Delta P$  [Pascals], the change in pressure as one ascends a distance  $\Delta z$  [m];  $\rho$  [kg m<sup>-3</sup>] the air density; g = 9.81 [m s<sup>-2</sup>] acceleration due to gravity.

•  $P = \rho R T$ 

The ideal gas law. P [Pascals], pressure;  $\rho$ , [kg m<sup>-3</sup>] the density; T [Kelvin], the temperature; and R = 287 [J kg<sup>-1</sup> K<sup>-1</sup>], the specific gas constant for air.

•  $e = \rho_v R_v T$ 

The ideal gas law for water vapor. e [Pascals], the vapour pressure (i.e. partial pressure of water vapour);  $\rho_v$ , [kg m<sup>-3</sup>] the absolute humidity (ie. vapor density); T [Kelvin], the temperature; and  $R_v = 462$  [J kg<sup>-1</sup> K<sup>-1</sup>], the specific gas constant for water vapor.

• 
$$q = \rho_v / \rho$$

the specific humidity

•  $L \uparrow = \epsilon \sigma T^4$ 

Stefan-Boltzmann law.  $L \uparrow [W m^{-2}]$ , the emitted longwave energy flux density;  $\epsilon$ , the emissivity of the surface (dimensionless);  $\sigma = 5.67 \times 10^{-8}$  [W m<sup>-2</sup> K<sup>-4</sup>], the Stefan-Boltzmann constant; T [K], the surface temperature.

• 
$$\frac{\Delta T}{\Delta z} = -0.01 \, [\mathrm{K} \, \mathrm{m}^{-1}]$$

The dry adiabatic lapse rate (DALR), i.e. for every one metre of ascent the temperature decreases by 0.01 degrees Kelvin.

•  $L_v = 2.5 \times 10^6 \, \mathrm{J \, kg^{-1}}$ 

The latent heat of vapourization of water

•  $\theta = 90 - \phi + \phi_{sol.dec}$ 

The solar elevation  $\theta$  at solar noon, at a location having latitude  $\phi$ , at the time of year when solar declination is  $\phi_{sol.dec}$ . Latitude is taken as positive for both hemispheres; solar declination is negative if the subsolar point is in the opposite hemisphere.