<u>Professor</u>: J.D. Wilson <u>Time available</u>: 45 mins <u>Value</u>: 20%

**Instructions**: For all 30 multi-choice questions, choose what you consider to be the best (or most logical) option. Use a pencil to mark that choice on the answer form. **Equations and data given at back**. **You may keep this exam** 

Note added after exam: in addition to understanding why the indicated answers are correct, it is equally useful to clarify for yourself why the other options would have been incorrect.

- 1. If a sample of air has vapour pressure e = 14 hPa then the dewpoint temperature of that sample  $T_d$  is closest to \_\_\_\_\_ °C
  - (a) 0
  - (b) 12  $\checkmark \checkmark$  [use Table]
  - (c) 14
  - (d) 18
  - (e) 20
- 2. If a sample of air has relative humidity RH = 50% and its temperature is  $T_d = 19^{\circ}$ C then its vapour pressure e is closest to \_\_\_\_\_ hPa
  - (a) 3
  - (b) 5
  - (c) 7
  - (d) 9

(e) 11 
$$\checkmark \checkmark$$
  $[e_s(19) = 22 \text{ hPa}; e = 22 \times 50/100 = 11 \text{ hPa}]$ 

- 3. The possibility to have formation of haze at relative humidities less than 100% owes to the presence in the air of \_\_\_\_\_
  - (a) condensation nuclei  $\checkmark \checkmark$  [p148, p164]
  - (b) ice nuclei
  - (c) supersaturation
  - (d) supercooled cloud droplets
  - (e) silver iodide
- 4. On Mars the gravitational acceleration  $g = 3.7 \,\mathrm{m \, s^{-2}}$ , and the Martian atmosphere of carbon dioxide has specific heat capacity  $c_p = 736 \,[\mathrm{J \, kg^{-1} \, K^{-1}}]$ . Thus on Mars a parcel lifted dry adiabatically through a distance of 100 m would \_\_\_\_\_
  - (a) warm by about 5 K
  - (b) cool by about 5 K
  - (c) warm by about 0.5 K
  - (d) cool by about 0.5 K  $\checkmark \checkmark$  [substitute into given eqn for lapse rate]
  - (e) cool by about 0.05 K

- 5. In adiabatic ascent of an unsaturated parcel, which water vapour variable remains unchanged?
  - (a) Relative humidity RH
  - (b) Absolute humidity  $\rho_v$
  - (c) Specific humidity  $q \checkmark \checkmark$  [p138]
  - (d) Vapour pressure e
  - (e) Dewpoint temperature  $T_d$

6. "Absolute instability" of the atmosphere commonly occurs

- (a) in the lower stratosphere
- (b) in the upper troposphere
- (c) at night
- (d) in a shallow ground-based layer during daytime heating  $(Q_H > 0) \checkmark \checkmark$  [p176; slide #6, lecture 16]
- (e) during the formation of a radiation fog

7. Collision efficiency for cloud droplets of radius r, R \_\_\_\_\_

- (a) is near unity for  $r \ll R$
- (b) is near unity for  $r \gg R$
- (c) is near unity for  $r \approx R$
- (d) is a maximum in warm clouds
- (e) is much smaller than unity if  $r \ll R$  or if  $r \gg R$   $\checkmark \checkmark$  [p205]

8. The radius of a typical cloud droplet is of order \_\_\_\_\_

[p203]

- (a)  $0.1 \ \mu m$
- (b) 10  $\mu$ m  $\checkmark \checkmark$  [Fig. 7-3]
- (c) 1000  $\mu m$
- (d) 10 mm
- (e) 10 cm
- 9. Small spherical particles of radius r fall relative to still air with a terminal velocity  $V_t$  that varies in proportion to \_\_\_\_\_
  - (a) 1/r (i.e. terminal velocity is halved if radius is doubled)

(b) 
$$1/\sqrt{r}$$
  
(c)  $\sqrt{r} \checkmark \checkmark$ 

- (d) r
- (e)  $r^2$

- 10. The Bergeron process for migration of water from supercooled droplets to ice crystals depends on the difference in \_\_\_\_\_ between surfaces of ice and water
  - (a) equilibrium vapor pressure  $\checkmark \checkmark$  [p206]
  - (b) temperature
  - (c) density
  - (d) terminal velocity
  - (e) vertical velocity
- 11. Specific humidity is the ratio  $q = m_v/m$  of the mass of water vapour in a given volume to the total mass  $m = m_v + m_d$ . If an unsaturated parcel is lifted \_\_\_\_\_
  - (a) q decreases at the Saturated Adiabatic Lapse Rate (SALR)
  - (b) q decreases at the Dry Adiabatic Lapse Rate (DALR)
  - (c) q decreases at the Environmental Lapse Rate (ELR)
  - (d) q increases
  - (e) q remains unchanged  $\checkmark \checkmark$  [p138]
- - (a) ten; entrainment of colder environmental air; more
  - (b) ten; entrainment of warmer environmental air; less
  - (c) ten; release of latent heat of condensation; less
  - (d) one; release of latent heat of condensation; more
  - (e) one; release of latent heat of condensation; less  $\checkmark \checkmark$  [p155]
- 13. The collection of supercooled cloud droplets by a falling ice crystal is a process called
  - (a) the solute effect
  - (b) the curvature effect
  - (c) riming  $\checkmark \checkmark$  [p207]
  - (d) aggregation
  - (e) graupellation
- 14. When a warm, almost saturated wind blows over a frozen lake, the result may be \_\_\_\_\_
  - (a) lake-effect snowfall
  - (b) advection fog  $\checkmark \checkmark$  [p160]
  - (c) formation of convective cloud over the lake
  - (d) freezing rain
  - (e) radiation fog

15. Which association is false?

- (a) Cirrus wispy
- (b) Cumulus heapy
- (c) Stratus layered
- (d) Nimbus producing rain, hail or snow
- (e) Orographic produced by dry, cool airflow over very warm ocean  $\checkmark$  [p170]
- 16. If the sky is overcast yet the sun casts shadows and is surrounded by a halo, the cloud type is \_\_\_\_\_
  - (a) Cirrostratus  $\checkmark \checkmark$  [p186]
  - (b) Sratocumulus
  - (c) Stratus
  - (d) Nimbostratus
  - (e) Altostratus
- 17. Suppose in a certain layer of the atmosphere the environmental lapse rate  $\text{ELR} = +0.05^{\circ}\text{C}\,\text{m}^{-1}$ , ie. for every 1 m increase in altitude, the temperature increases by 0.05°C. This layer is
  - (a) unconditionally stable  $\checkmark \checkmark$  [p178; it was noted in the exam that "unconditionally" and "absolutely" have the same meaning]
  - (b) unconditionally unstable
  - (c) conditionally unstable
  - (d) conditionally stable
  - (e) neutral with respect to dry adiabatic motion
- 18. The idealized three-cell model of the general circulation predicts low-level easterlies at low latitude ('trade winds') and at high latitude ('polar easterlies'). In reality easterly trade winds are \_\_\_\_\_ and the polar easterlies are \_\_\_\_\_
  - (a) a persistent weather feature; a persistent weather feature
  - (b) a persistent weather feature; visible only as a climatological feature  $\sqrt{\sqrt{}}$  [p231]
  - (c) visible only as a climatological feature; a persistent weather feature
  - (d) westerlies; easterlies
  - (e) easterlies; westerlies
- 19. The 3-cell model for the General Circulation suggests that at latitude 30 degrees one will find \_\_\_\_\_ surface pressure while at latitude 60 degrees one will find \_\_\_\_\_ surface pressure.
  - (a) Low; low
  - (b) High; high
  - (c) High; low  $\checkmark \checkmark$  [Fig. 8-3]
  - (d) Low; high
  - (e) Negative; positive

- 20. Regarding semipermanent pressure systems of the general circulation, that which is most significant for winter weather in western Canada is the \_\_\_\_\_
  - (a) Siberian High
  - (b) Aleutian Low  $\checkmark \checkmark$  [p234; slide #8, lec19]
  - (c) Icelandic Low
  - (d) Hawaiian High
  - (e) Tibetan Low
- 21. Strong deviations of temperature away from the climatological norm are likely to occur during a sustained period of \_\_\_\_\_ flow associated with a \_\_\_\_\_ long wave (Rossby wave) pattern
  - (a) zonal; large amplitude
  - (b) zonal; gently meandering
  - (c) meridional; large amplitude  $\checkmark \checkmark$  [p239]
  - (d) meridional; gently meandering
  - (e) westerly; gently meandering

22. An afternoon sea-breeze is most likely to occur

- (a) after a sunny, calm morning  $\checkmark \checkmark$
- (b) after a cloudy, windy morning
- (c) after a cloudy, calm morning
- (d) after a rainy, calm morning
- (e) in conjunction with a cyclonic storm
- 23. The Walker circulation is a feature of the observed General Circulation (GC) in the equatorial Pacific that \_\_\_\_\_
  - (a) redistributes energy and water vapour across longitude lines  $\checkmark \checkmark$  [Fig. 8-29]

[maximizing the differential heating rate]

- (b) redistributes energy and water vapour across latitude lines
- (c) is one of the three "cells" of the 3-cell GC Model
- (d) occurs within the polar cell of the 3-cell GC Model
- (e) is unrelated to the El-Nino phenomenon
- 24. A deep layer of fog is more likely to form on a night with a very gentle wind, than during a night which is completely calm, because \_\_\_\_\_
  - (a) rate of longwave radiant emission from ground is increased by wind
  - (b) eddies carry heat down to the cooling surface from a deeper layer  $\checkmark \checkmark$  [p159]
  - (c) a light wind increases the rate of cooling of the ground
  - (d) the ELR equals the DALR
  - (e) the light breeze prevents gravitational settling of condensation nuclei

## For the remaining questions, please refer to the attached charts.

- 25. Heavy short-dashed lines on the skew T-log p diagram (Figure 1) identify several families of reference curves. The family of dry adiabats is represented by line \_\_\_\_\_
  - (a) A
  - (b) B  $\checkmark \checkmark$  [slide #11 lec13]
  - (c) C
  - (d) D
  - (e) E

## 26. Layer L1 should be classified as \_\_\_\_\_

- (a) absolutely unstable
- (b) absolutely stable
- (c) conditionally unstable
- (d) conditionally stable
- (e) neutral with respect to dry adiabatic motion  $\sqrt{\checkmark}$  [slide #20, lec14]

## 27. The ground-based layer below layer L1 is \_\_\_\_\_

- (a) absolutely unstable  $\checkmark \checkmark$
- (b) absolutely stable
- (c) conditionally unstable
- (d) conditionally stable
- (e) neutral with respect to dry adiabatic motion

28. Comparing Figures (2,3) the change in 1000-500 hPa thickness at Edmonton was about

- (a) 35 m
- (b) 62 m
- (c) 35 dam  $\checkmark \checkmark$  [thickness changes from about 549 dam to about 514 dam]
- (d) 62 dam
- (e) 350 dam
- 29. Corresponding to the above thickness change of n dam, the mean temperature of the 1000-500 hPa layer over Edmonton had \_\_\_\_\_ during the interval by \_\_\_\_\_ degrees Kelvin
  - (a) warmed; 2n
  - (b) cooled; 2n
  - (c) warmed; n/2
  - (d) cooled;  $n/2 \checkmark \checkmark$  [slide #2 lec15]

- 30. Figure (4) gives the Edmonton sounding data up to 700 hPa for two times. The change in thickness of the 925-850 hPa layer during the interval was a/an \_\_\_\_\_ of about \_\_\_\_\_ [m]
  - (a) decrease; 40  $\checkmark \checkmark$  [slide #10 lec15. At the earlier time: 1488-792=696 m; at later time 1416=7-761=656 m. Thus a decrease of 40 m]
  - (b) decrease; 700
  - (c) increase; 40
  - (d) increase; 700

## Equations and Data

$$\bullet \ Q^* = Q_H + Q_E + Q_G$$

Surface energy balance on a reference plane at the base of the atmosphere, all fluxes in  $[W m^{-2}]$ .  $Q^*$  the net radiation, positive if directed towards the surface;  $Q_H, Q_E$  the sensible and the latent heat fluxes, positive if directed from the surface towards the atmosphere;  $Q_G$  the 'soil' heat flux, positive if directed from the surface into ground/lake/ocean.

•  $\frac{\Delta T}{\Delta z} = -\frac{g}{c_p}$ 

The dry adiabatic lapse rate, where g is the gravitational acceleration and  $c_p [J \text{ kg}^{-1} \text{ K}^{-1}]$  is the specific heat at constant pressure.

•  $e, \rho_v, q = m_v/(m_v + m_d) = \rho_v/\rho, T_d, \text{RH} = 100 \, e/e_s(T)$ 

Humidity variables: e [Pa] the vapour pressure;  $\rho_v$  [kg m<sup>-3</sup>] the absolute humidity vapour density); q the specific humidity (where  $m_v$ ,  $m_d$  are respectively the mass of vapour and of "dry air" in a sample);  $T_d$  the dewpoint temperature; and RH the relative humidity (where  $e_s(T)$  is the equilibrium vapour pressure corresponding to temperature T of the sample)

Table 1: Equilibrium vapour pressure  $e_s(T)$  [hPa] versus temperature T [°C]. Figure cited applies to equilibrium over a plane surface of water where  $T \ge 0^{\circ}$ C, or of ice where  $T < 0^{\circ}$ C.

T	$e_s(T)$	T	$e_s(T)$	T	$e_s(T)$	T	$e_s(T)$	T	$e_s(T)$	T	$e_s(T)$	T	$e_s(T)$
-5	4.02	0	6.11	5	8.72	10	12.27	15	17.04	20	23.37	25	31.67
-4	4.37	1	6.57	6	9.35	11	13.12	16	18.17	21	24.86	26	33.61
-3	4.76	2	7.05	7	10.01	12	14.02	17	19.37	22	26.43	27	35.65
-2	5.17	3	7.58	8	10.72	13	14.97	18	20.63	23	28.09	28	37.80
-1	5.62	4	8.13	9	11.47	14	15.98	19	21.96	24	29.83	29	40.06



Figure 1: Thermodynamic chart for Edmonton, 00Z 1 Oct. 2009.



Figure 2: CMC 500 hPa analysis for 00Z 21 April 2009. Boxes with white text on a black background label height contours, boxes with black text on a white background label thickness contours.



Figure 3: CMC 500 hPa analysis for 12Z 24 April 2009. Boxes with white text on a black background label height contours, boxes with black text on a white background label thickness contours.

00Z,	21	April	2009
		123772	

12Z, 24 April 2009

\_\_\_\_\_

PRES hPa	HGHT m	TEMP C	PRES hPa	HGHT m	TEMP C		
1000.0 928.0 926.0 925.0 911.4 878.2 850.0 846.2 836.0 815.1 807.0 784.9 781.0 784.9 781.0 755.6 727.2 700.0	134 766 783 792 914 1219 1488 1524 1623 1829 1910 2134 2174 2438 2743 3047	12.2 11.4 11.2 10.1 7.4 5.0 4.6 3.6 2.2 1.9 1.8 -0.3 -2.7 -5.1	1000.0 925.0 924.0 919.0 906.7 887.0 872.0 850.0 850.0 838.1 805.3 792.0 773.5 755.0 742.6 712.7 708.0 700.0		-6.7 -6.9 -7.4 -8.3 -9.4 -11.1 -11.9 -14.3 -15.3 -16.7 -18.1 -19.1 -21.7 -22.1 -22.5		

Figure 4: Edmonton soundings.