

Chapter 1 sets the stage and addresses such basics as:

- What are the most important components of the atmosphere, and why?
- How much mass is in one cubic metre of “air” ? How much oxygen?
- Does the air get colder or warmer as you ascend from ground?
- What is the difference between weather and climate?
- Why is pressure so important to meteorologists?
- Where in the atmosphere does weather “happen”?

Note that the aim of the course – and of the authors of our textbook (see p4) – is to help you “understand the underlying physical processes.” The challenge is to develop one's own judgment as to *what is most important* – you are not expected to remember every detail and every digit...

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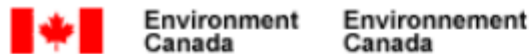
***Here's the scope of our course***

# “Weather” versus “climate” (p4)

Weather: short-term, local phenomena

Climate: long term patterns (statistics: averages, extremes, frequencies of events...)

“Climate is what you expect but weather is what you get”




[Home](#) > [Current Conditions and Forecasts](#) > [Alberta](#) >

[http://www.weatheroffice.gc.ca/canada\\_e.html](http://www.weatheroffice.gc.ca/canada_e.html)

## Edmonton

Z (“Zulu”) = GMT = UTC

### Current Conditions

 <b>30 °C</b>	Observed at:	<b>Edmonton City Centre Airport</b>		
	Date:	<b>4:00 PM MDT Thursday 8 September 2011</b>		
	Condition:	<b>Sunny</b>	Temperature:	<b>29.8°C</b>
	Pressure:	<b>102.1 kPa</b>	Dewpoint:	<b>9.7°C</b>
	Tendency:	<b>falling</b>	Humidity:	<b>29 %</b>
	Visibility:	<b>15 km</b>	Wind:	<b>SSE 9 km/h</b>
	<a href="#">Air Quality</a>	<b>5</b>	<a href="#">Humidex:</a>	<b>31</b>
<a href="#">Health Index:</a>				

**MDT = GMT - 6**

### Historical Data

Yesterday		Normals	
Max:	<b>29.8°C</b>	Max:	<b>18°C</b>
Min:	<b>10.7°C</b>	Min:	<b>5°C</b>
<a href="#">Precip:</a>	<b>0.5 mm</b>		

“It is normal for the weather to be abnormal”

# Vertical Structure – and thinness of atmosphere

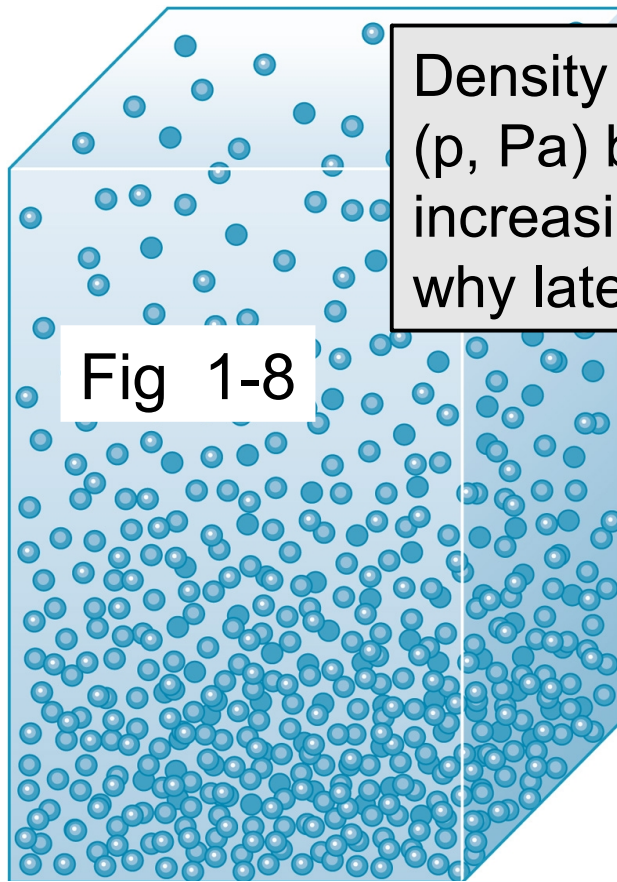


Fig 1-8

Density ( $\rho$ ,  $\text{kg m}^{-3}$ ) and pressure ( $p$ , Pa) both decrease with increasing elevation – shall see why later, Ch 4

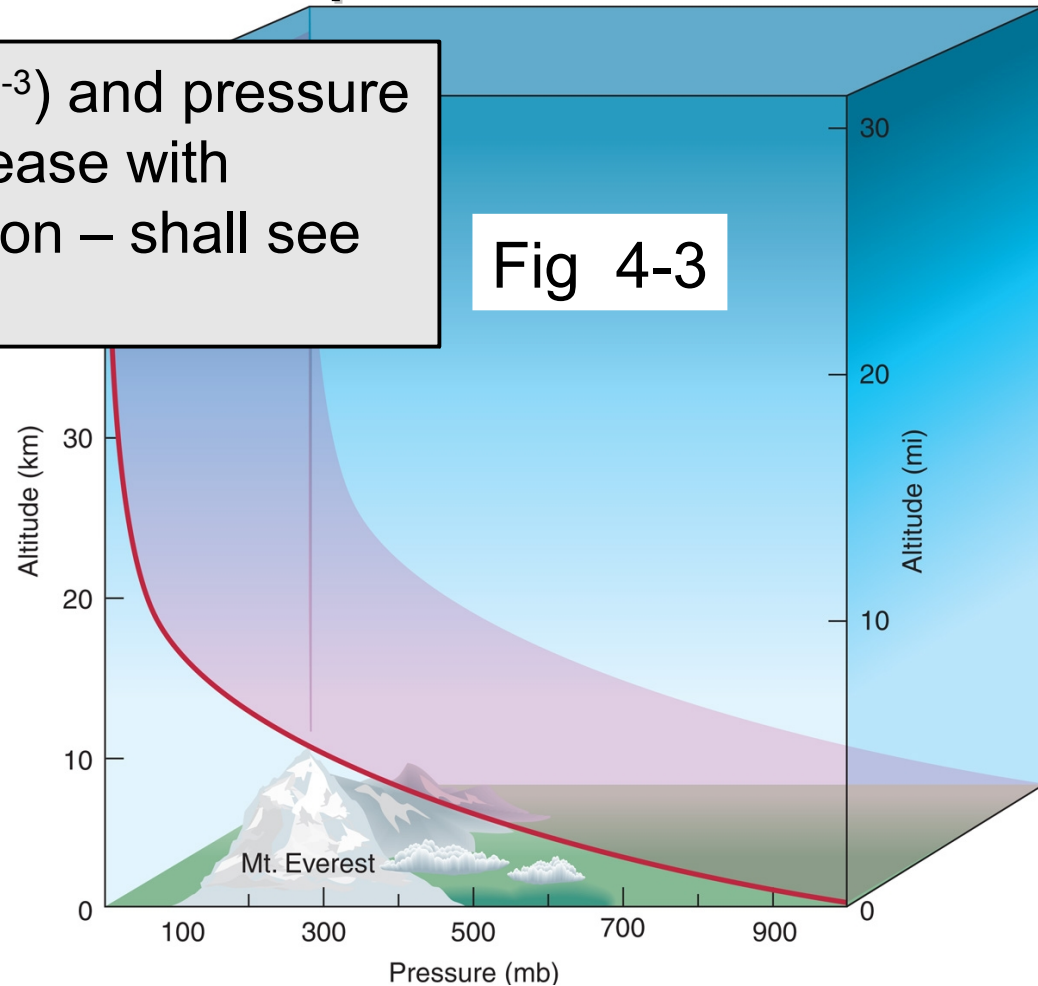


Fig 4-3

“Given the shallowness of the atmosphere, its motion (*when averaged*) over large areas must be primarily horizontal”... yet “Paradoxically, the least impressive motions – vertical – determine much of atmospheric behaviour”

# “Mandatory” (Conventional) Reporting/Analysis Levels

Increasing height, decreasing pressure



250 hPa ,  $z \approx 10.5 \text{ km} = 1050 \text{ dam}$   
(34,000 feet)

500 hPa ,  $z \approx 5.5 \text{ km} = 550 \text{ dam}$

700 hPa ,  $z \approx 3 \text{ km} = 300 \text{ dam}$

850 hPa ,  $z \approx 1.5 \text{ km} = 150 \text{ dam}$

Edmonton Int'l,  $z = 723 \text{ m}$

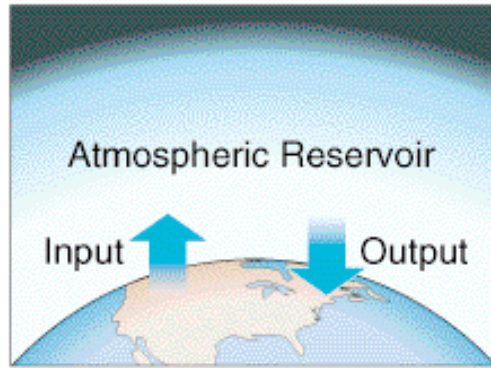
MSL:  $p \approx 1000 \text{ hPa}$ ,  $z \equiv 0 \text{ m}$

density  $\rho \approx 1 \text{ kg m}^{-3}$

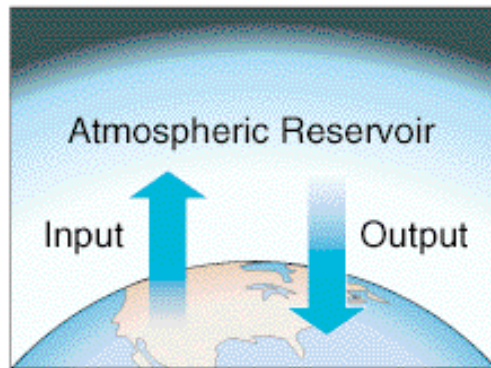
1 millibar (mb) = 1 hectoPascal (hPa) = 100 Pascals (Pa)

# Atmospheric composition

Fig. 1-2



(a)



(b)

uniformly-mixed in the “homosphere” (lowest 80 km) and “inactive”

## Permanent Gases

- N<sub>2</sub> 78 % (by volume)
- O<sub>2</sub> 21 %
- Ar 1 %

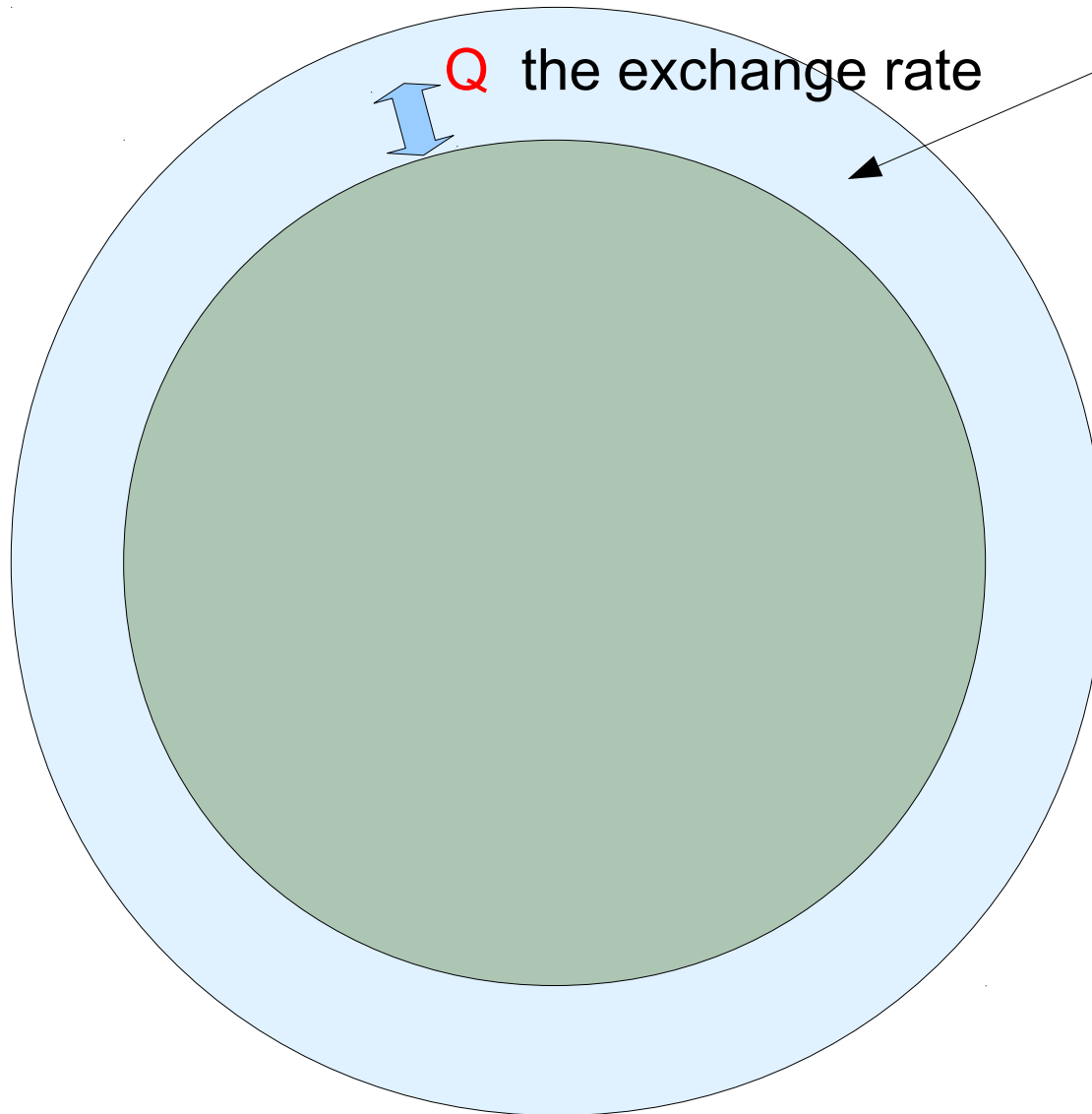
not uniformly-mixed, these gases “active”

## Variable Gases

- H<sub>2</sub>O < 4 % (1/4 % globally)
- CO<sub>2</sub> 0.038% (386 ppm)
- ozone
- methane
- 

Sfc-atmos. exchange by phys., chem. or biological processes + internal product'n/destruct'n

# Concept of “residence time”



**M** the total mass of a given gas averaged over a long time (years or longer)

- If the net transport **Q** from sfc to atmos vanishes over the averaging time and there is no net production internally, we say the concentration is in “steady state”

$N_2$  – a long residence-time gas (slow volcanic source) that is unimportant (inactive) with respect to “weather”

- crucial to plant nutrition
  - added to atmos by decay of organic material
  - removed by soil bacteria
- } “slowly” (small  $Q$ )

huge

$$T^{\text{res}} = \frac{M}{Q} \left[ \frac{\text{kg}}{\text{kg yr}^{-1}} \right] = \text{tens of millions of years}$$

small



H<sub>2</sub>O – a variable gas that is **not** uniformly mixed and is “active” (“thermal”)

- interacts with longwave radiation (“radiatively active” or “greenhouse” gas)
- affects atmospheric motion by means of buoyancy (moist air is lighter)
- every 1 kg of vapour stores 2.5 million Joules of latent heat (ie. energy)

small

$$T^{\text{res}} = \frac{M}{Q} \left[ \frac{\text{kg}}{\text{kg yr}^{-1}} \right] = \text{about 10 days}$$

large

(evaporation, precipitation)

To get a sense of the enormous amount of energy tied up in water vapour:

Calculation: how long ( “ $t$  ” ) must you run a 100 W light bulb to consume 2.5 million Joules of energy (i.e. an amount sufficient to evaporate 1 kg of water)?

Given, energy  $E$ :  $E = 2.5 \times 10^6 = 2,500,000 [J]$

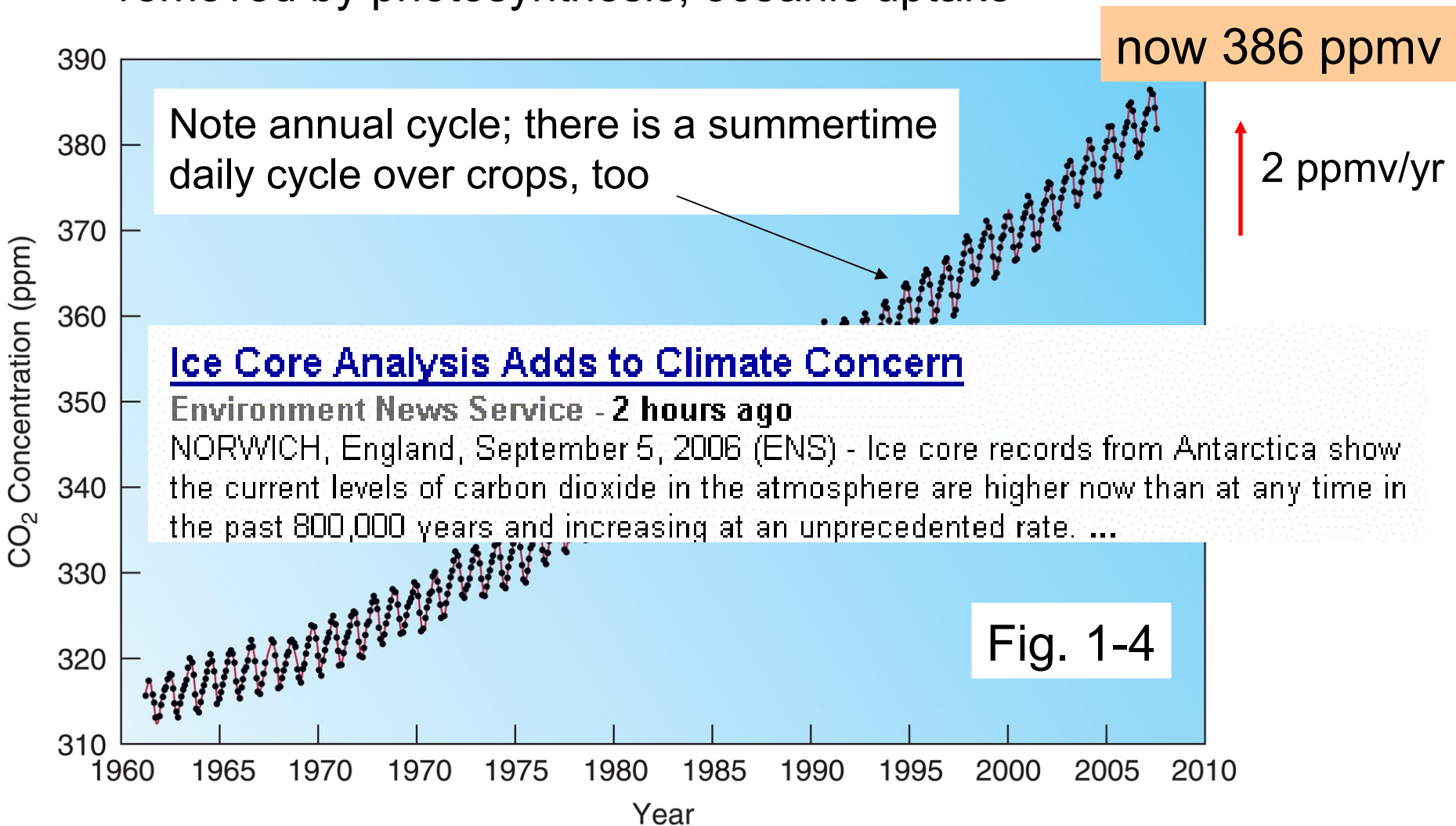
Formula:  $E = P t$

(where  $P$  = Power [W],  $t$  = time [s] )

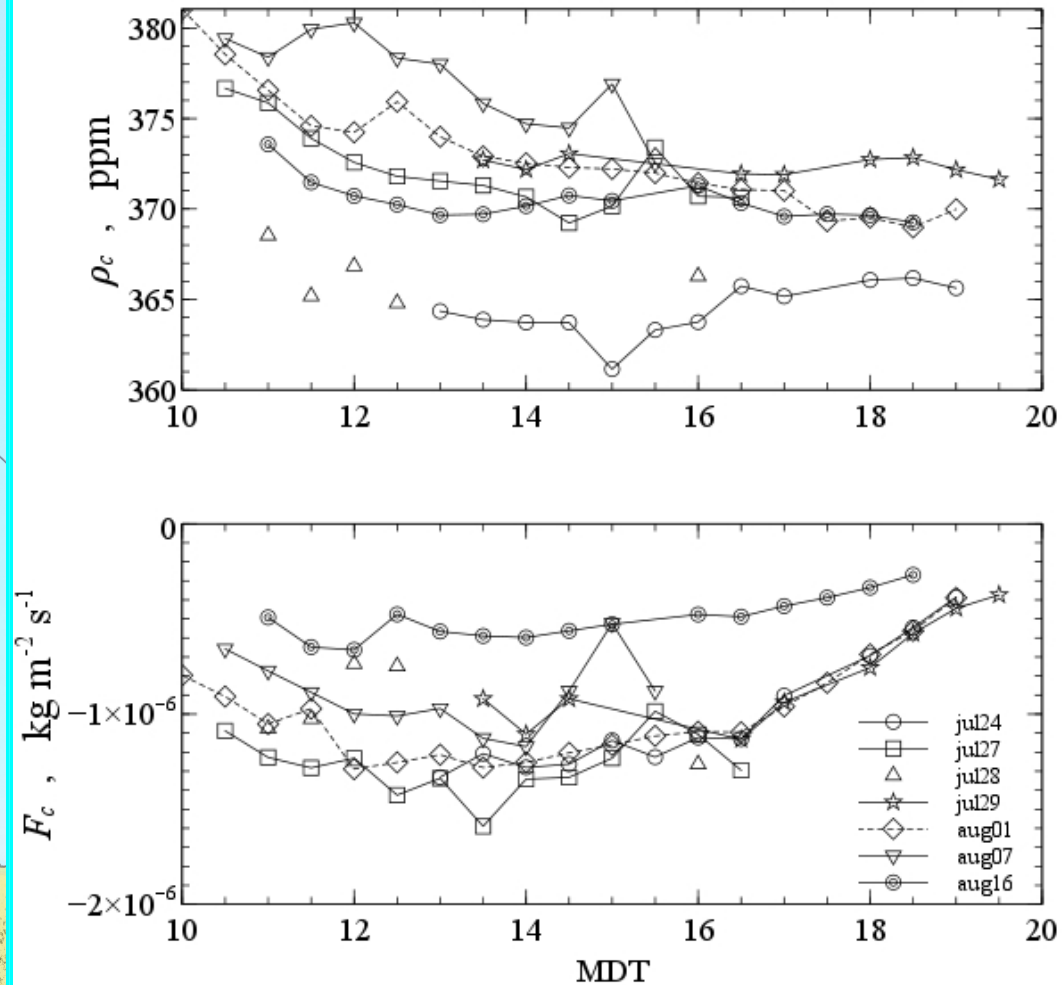
Unknown, time  $t$  :

CO<sub>2</sub> – a variable gas that is (fairly) uniformly mixed and is “active”

- radiatively active (greenhouse gas)
- added to atmos by volcanoes, respiration and decay of vegetation, breathing, burning fuels
- removed by photosynthesis, oceanic uptake



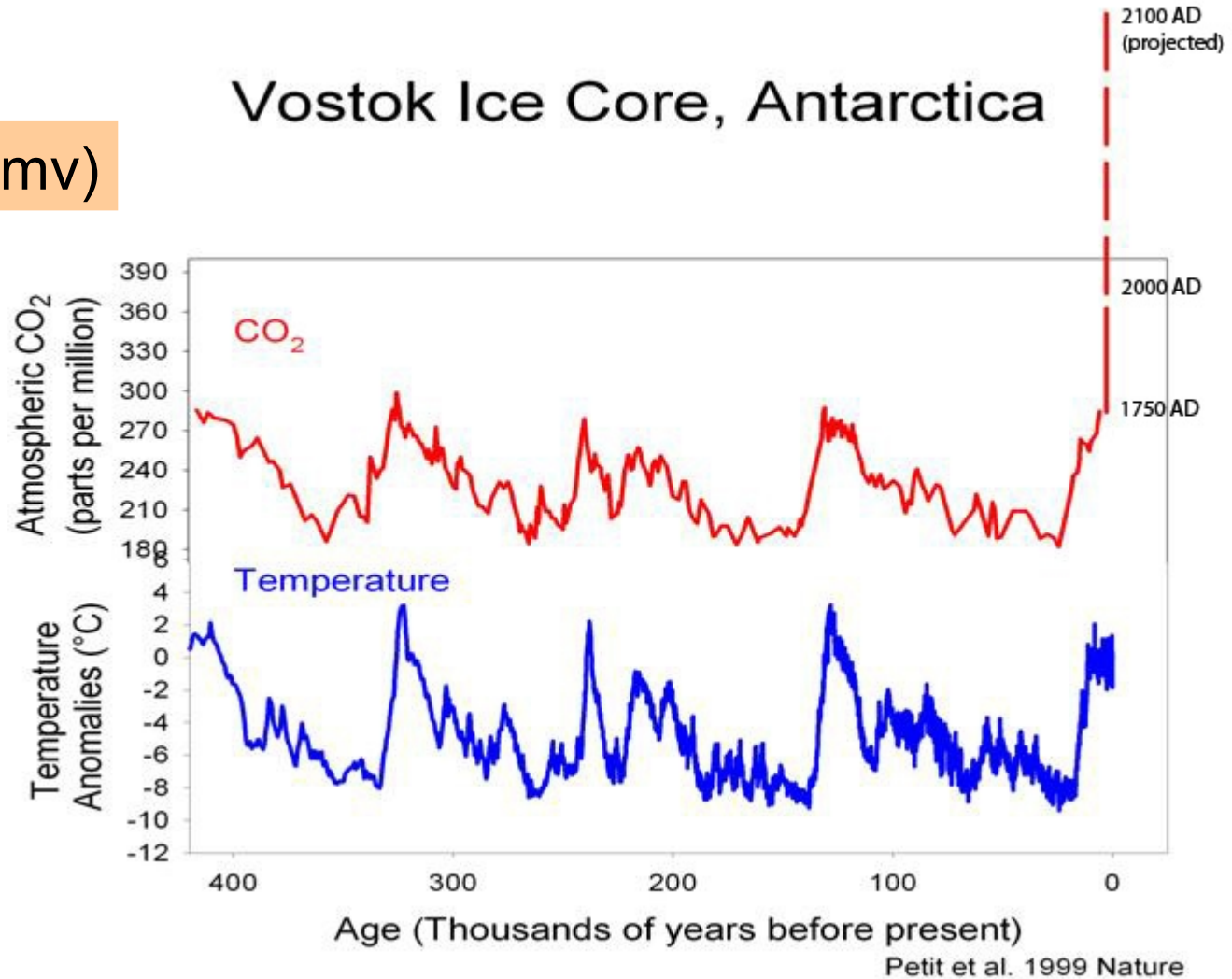
# Measured concentration and flux of CO<sub>2</sub> at St. Albert, 2011



- Wheat canopy is absorbing about \_\_\_ mg per square metre per hour

(now 386 ppmv)

## Vostok Ice Core, Antarctica



Dr Eric Wolff, British Antarctic Survey: “Ice cores reveal the Earth's natural climate rhythm over the last 800,000 years. When carbon dioxide changed there was always an accompanying climate change...”