<u>Professor</u>: J.D. Wilson <u>Time available</u>: 80 mins <u>Value</u>: 15%

#### Please answer in the booklet provided. Equations and data given at back.

## A. "Live" web weather data $(8 \ge 1/2 \rightarrow 4\%)$

To answer the following questions, please use whatever web resources best suit (e.g. CMC analyses, text data for Stony Plain sounding, etc).

- 1. Red Deer Regional Airport reported [minus 30 Celcius] for the 0600 MST temperature this morning
- 2. 1000-500 hPa thickness at Edmonton at 12Z this morning was [5143 m]
- 3. Use Vizaweb to access the 0h panel of the GEM Regional Run initialized at 12Z today; choose N\_America as Domain. The strongest updraft at the 850 hPa level has a magnitude in the range [-4 to -8] Pa s<sup>-1</sup>
- 4. According to the GEM Regional Model run initialized at 12Z today, the 1000-500 hPa thickness over Edmonton at 12Z Thurs 10 Feb. will be [about 526 528 dam, depending on which specific prog you used]
- 5. According to this morning's Edmonton sounding, true surface pressure was [943 hPa]
- 6. According to this morning's Edmonton sounding, height (AGL) of the 500 hPa surface over Edmonton was [5450-766=4684 m]
- 7. According to this morning's sounding from Fort Smith (YSM), temperature-dewpoint spread at 700 hPa was [34°C]
- 8. What WATCHES, WARNINGS, OR STATEMENTS are listed by the most recent SIG-NIFICANT WEATHER DISCUSSION ISSUED BY THE PRAIRIE AND ARCTIC STORM PREDICTION CENTRE OF ENVIRONMENT CANADA? [BLIZZARD WARNINGS FOR CAMBRIDGE BAY, GJOA HAVEN. BLIZZARD WARNING FOR MOST OF THE KIVALLIQ DISTRICT AND CLYDE RIVER. WIND CHILL WARNINGS FOR RESO-LUTE AND TALOYOAK. WIND CHILL WARNING FOR REGINA AND ESTEVAN.]

#### B. Interpretation of weather charts. $(6 \ge 1/2 \rightarrow 3\%)$

In Edmonton on January 15 the minimum and maximum temperatures were  $(-26.1, -22.0)^{\circ}$ C, while on Jan 20 they were  $(-14.9, +2.6)^{\circ}$ C. The weather situation on these two days is depicted by Figures (1, 2). Edmonton experienced snowfall on both days: 4.4 cm on the 15th, and 0.9 cm on the 20th.

- 1. Temperatures at the 850 hPa level over east-central Alberta were more uniform on \_\_\_\_\_ while low-level winds over SW Alberta would have been stronger on \_\_\_\_\_
  - (a) 15 Jan; 15 Jan
  - (b) 20 Jan; 20 Jan
  - (c) 15 Jan; 20 Jan  $\checkmark \checkmark$
  - (d) 20 Jan; 15 Jan
- The system far north of Manitoba (and visible on both days) is often termed the/an \_\_\_\_\_\_.
  The Alberta lee trough features on \_\_\_\_\_\_
  - (a) Arctic vortex; 15 Jan
  - (b) Alberta clipper; 15 Jan
  - (c) Arctic vortex; 20 Jan  $\checkmark\checkmark$
  - (d) Alberta clipper; 20 Jan
- 3. Lloydminister (east of Edmonton, on the Saskatchewan border) should have seen \_\_\_\_\_\_\_\_ on \_\_\_\_\_\_
  - (a) steady warming; 20 Jan  $\checkmark \checkmark$
  - (b) steady cooling; 20 Jan
  - (c) strong low level winds; 15 Jan
  - (d) mild surface temperatures; 15 Jan
- 4. On neither day was wind data submitted from the sounding at Fort Smith<sup>1</sup> (YSM, 71934; just north of the N. border of Alberta). In the framework of the Geostrophic model, a reasonable categorical guess for the Fort Smith 700 hPa winds on 15th and 20th Jan. (respectively) would be
  - (a) calm; moderate WNW
  - (b) calm; moderate ESE
  - (c) weak to moderate from sector  $135 225^{\circ}$ ; moderate WNW  $\checkmark \checkmark$
  - (d) weak to moderate from sector  $315 360^{\circ}$ ; moderate WNW

<sup>&</sup>lt;sup>1</sup>or at The Pass (YQD, 71867; near the Saskatchewan-Manitoba border).

- 5. Comparing Edmonton's 850 and 700 hPa winds on 20 Jan., we note "veering," which is associated warm advection. The atmosphere over Central Alberta at this time was \_\_\_\_\_\_ and the thermal wind  $\vec{U}_{700|850}$  was probably
  - (a) baroclinic; oriented roughly parallel to the 850 hPa height contours
  - (b) baroclinic; oriented roughly parallel to the 850 hPa isotherms  $\checkmark\checkmark$
  - (c) barotropic; oriented roughly perpendicular to the 850 hPa isotherms
  - (d) barotropic; oriented roughly parallel to the 850 hPa isotherms
- 6. On both occasions there was strong, moist WNW 700 hPa flow over central Alberta, and overall the 700 hPa flow patterns on these two days appear qualitatively similar; *surface* conditions, however, were very different. Which of the following statements makes most sense?
  - (a) the flow at the 700 hPa level *exerts no influence*, direct or indirect, on surface temperature
  - (b) sometimes indeed, frequently present conditions cannot be explained in terms of (or attributed to) the *present* configuration of the circulation pattern, for present conditions depend (also) on the *recent history* of the circulation, at all levels  $\checkmark \checkmark$
  - (c) the observed very different surface conditions have to be explainable in terms of more subtle elements of the upper flow, such as diffluence (Jan 15) or a weak shortwave (Jan 20) over north-central Alberta
  - (d) the greater snowfall amount on 15 Jan is a direct consequence of the much lower surface temperature, which resulted in a reduced upward longwave radiative energy flux density, cooling the clouds

#### C. Calculations & use of thermodynamic chart $(4 \ge 2 \ge 8 \%)$

Morning temperatures at Edmonton City Centre Airport hit  $-30^{\circ}$ C on Monday 31 Jan 2011. Figure (3) is the sounding and Fig. (4) the 850 hPa analysis for 12Z on Tues 1 Feb.

- 1. Referring to Figure (3), if a parcel "P" of air from 700 hPa were lowered adiabatically to 925 hPa, what would be the resulting values of its temperature and dewpoint? If the parcel were instead *lifted* adiabatically, its temperature-dewpoint spread would converge to zero at what pressure level and at what temperature? Answers: by tracing a dry adiabat down to 925 hPa one finds  $T \approx +11^{\circ}$ , and by tracing a mixing-ratio line (same as dewpoint lapse rate line) down one has  $T_d \approx -28^{\circ}$ . Conversely, moving upward from 700 hPa the dry adiabat and the mixing-rtio line intersect at about 500 hPa with  $T = T_d \approx -36^{\circ}$ .
- 2. What would the parcel P's density, vapour pressure, absolute humidity and specific humidity be, when lowered to 925 hPa? Answers:  $\rho \approx 1.13 \text{ kg m}^{-3}$ ; and with  $e = e_*(-28) \approx 47$ Pa we get  $\rho_v \approx 3.6 \times 10^{-4} \text{ kg m}^{-3}$ ; thus  $q \approx 3.2 \times 10^{-4} \text{ kg kg}^{-1}$
- 3. Referring to Fig. (4), compute an approximate value for the Geostrophic 850 hPa windspeed at the point marked "X" in NW Saskatchewan. *Answer*: Latitude is about 58°. The distance between neightbouring height contours is  $\Delta n \approx (10.5 \text{ mm})/46 \text{ mm}) \times 11 \times 111,000$ m (about 280 km).  $V \approx 17 \text{ m s}^{-1}$ .

4. Again referring to Fig. (4) and using the wind speed computed above, calculate an approximate value for the rate of temperature advection  $A_T$  at the point marked "X." Please give your answer in °C hr<sup>-1</sup>, and state whether your result corresponds to warming or cooling. *Answer*: Roughly +1°C/hr, warming.

### Equations and Data.

- one full barb on the wind vector corresponds to 5 m s<sup>-1</sup>, and 1 degree of latitude corresponds to a distance of 111 km
- $p = \rho R T$ , the ideal gas law. p [Pascals], pressure;  $\rho$ , [kg m<sup>-3</sup>] the density; T [Kelvin], the temperature; and R = 287 [J kg<sup>-1</sup> K<sup>-1</sup>], the specific gas constant for air.
- $e = \rho_v R_v T$ , the ideal gas law for water vapour. e [Pascals], vapour pressure;  $\rho_v$ , [kg m<sup>-3</sup>] the absolute density; T [Kelvin], the temperature; and  $R_v = 462$  [J kg<sup>-1</sup> K<sup>-1</sup>], the specific gas constant for water vapour.
- $\theta = T\left(\frac{p_0}{p}\right)^{R/c_p}$ , the potential temperature  $\theta$  [K] of air whose actual pressure and temperature are (p, T), i.e. the temperature that air would have if compressed adiabatically to pressure  $p_0$ . The exponent involves the gas constant for air  $(R = 287 \text{ J kg}^{-1} \text{ K}^{-1})$  and the specific heat of air at constant pressure  $(c_p \approx 1000 \text{ J kg}^{-1} \text{ K}^{-1})$ . Temperatures must be expressed in the Kelvin unit.
- the saturation vapour pressure at temperature T is given by

$$e_*(T) \approx 610.78 \exp \frac{19.8 T}{273 + T}$$
 (over water),  
 $\approx 610.78 \exp \frac{22.5 T}{273 + T}$  (over ice),

where T is to be entered in degrees Celcius and e is in [Pa].

• 
$$A_T \equiv \left(\frac{\partial T}{\partial t}\right)_{adv} = -V \frac{\partial T}{\partial s}$$

Advective contribution to the rate of change of temperature, expressed in natural coordinates. The unit vector  $\hat{s}$  for the *s* axis points downstream and parallel to the flow contours (eg. height contours), and *V* is the wind *speed*.

• 
$$V = \frac{g}{f} \frac{\Delta h}{\Delta n}$$

The Geostrophic wind equation.  $\Delta h$  [m], the change in height of a constant pressure surface over distance  $\Delta n$  [m] normal to the height contours;  $f = 2\Omega \sin \phi$  [s<sup>-1</sup>] the Coriolis parameter (where  $\Omega \approx 2\pi/(24 \times 3600)$  s<sup>-1</sup> is the angular velocity of the earth, and  $\phi$  is latitude); g acceleration due to gravity.

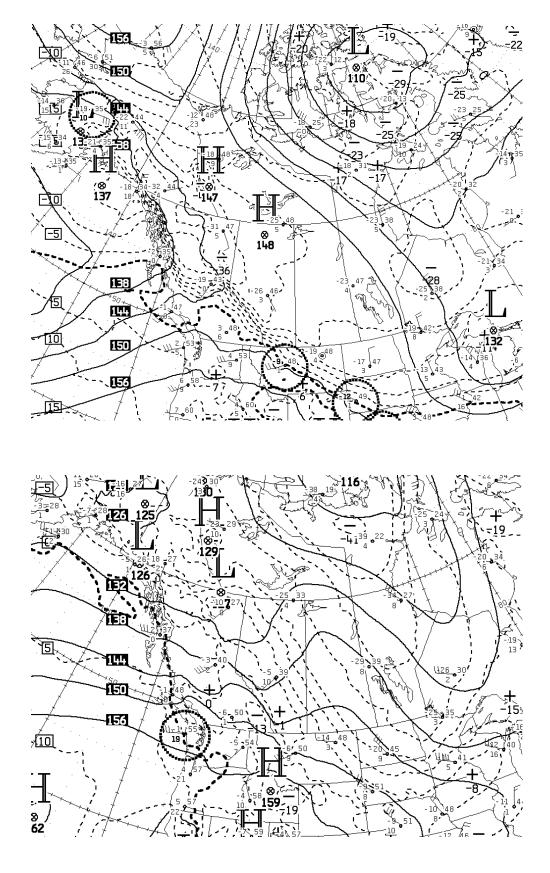


Figure 1: CMC 850 hPa analyses for 12Z on January 15th (upper) and 20th (lower), 2011.

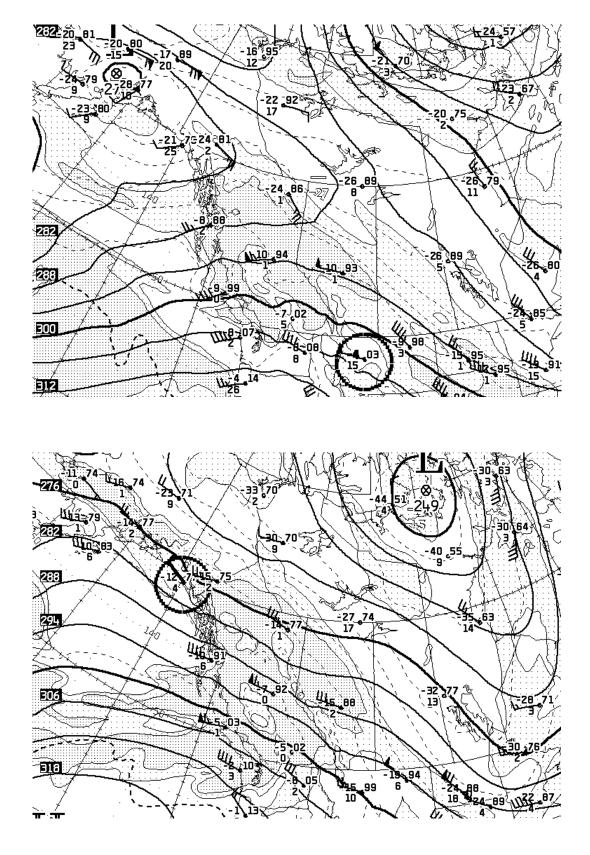


Figure 2: CMC 700 hPa analyses for 12Z on January 15th (upper) and 20th (lower), 2011.

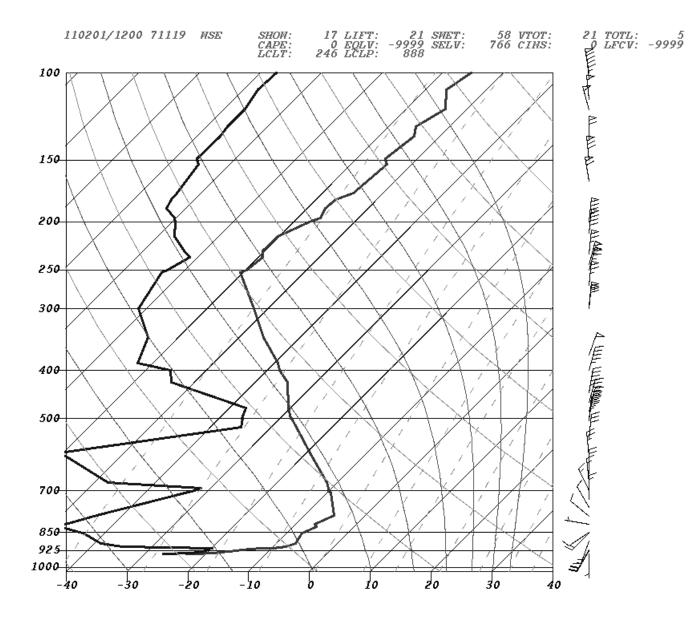


Figure 3: Edmonton (Stony Plain) sounding at 12Z on 1 Feb., 2011.

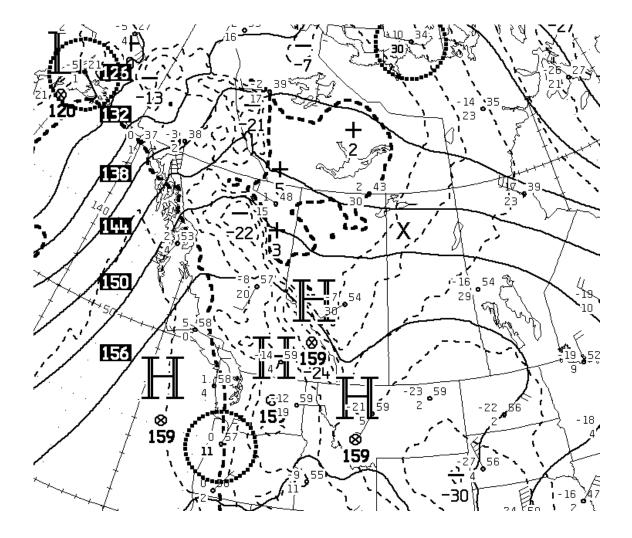


Figure 4: CMC 850 hPa analysis, 12Z on 1 Feb., 2011.

# Class structure, environment, delivery

— your (anonymous) feedback will be appreciated